

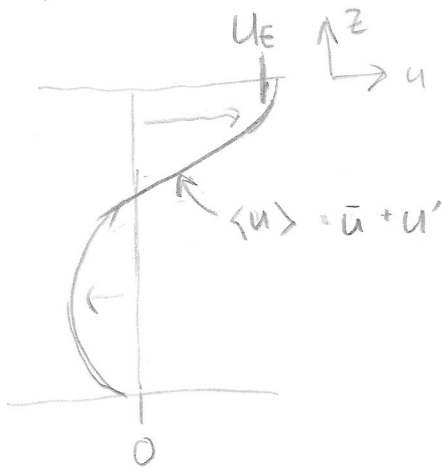
Estuarine Salinity Structure

tidally-averaged

8/7/2019

(1)

Recall: the exchange flow looks like:



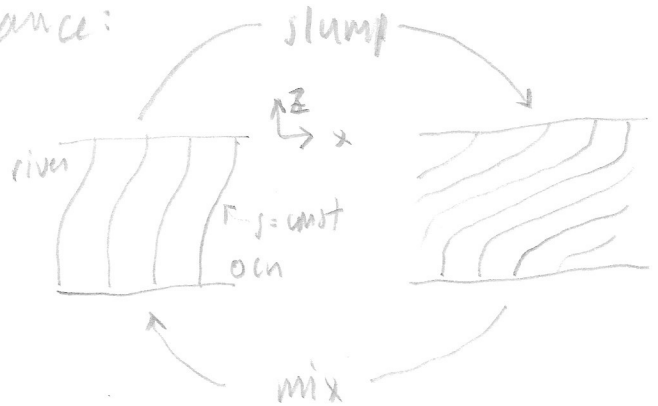
$$\text{where } U_E = \frac{g\beta\bar{s}_x H}{4g} \frac{H^2}{A}$$

represents the balance between pressure gradient (increases exchange) and friction (slows exchange)

The stratification $s'(z)$ is mainly controlled by a similar local balance:

$$s'_t = -u' \bar{s}_x + K s'_{zz}$$

↑ slump
↑ mix



Find steady solution: $s'_{zz} = \frac{\bar{s}_x}{K} u'$ with $s'_z = 0$ at $z = -H, 0$
 (no flux through bottom or top)

Result

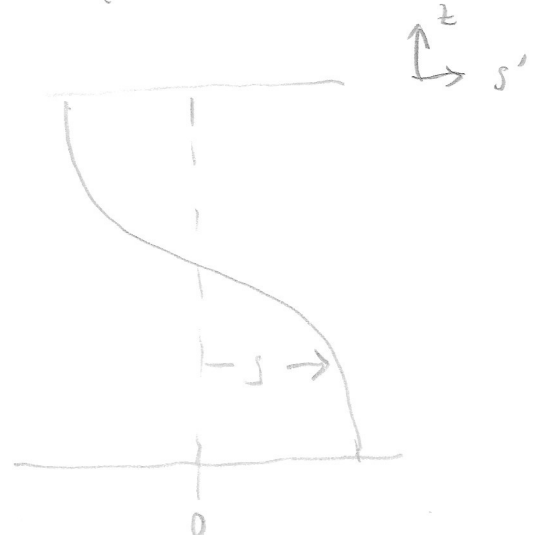
$$s' = \frac{H^2}{K} \bar{s}_x (\bar{u} P_1 + U_E P_2)$$

where $P_1 = -\frac{1}{120} + \frac{1}{4} \zeta^2 - \frac{1}{8} \zeta^4$

$$P_2 = -\frac{1}{12} + \frac{1}{2} \zeta^2 - \frac{3}{4} \zeta^4 - \frac{2}{5} \zeta^5$$

and $\zeta = \frac{z}{H}$

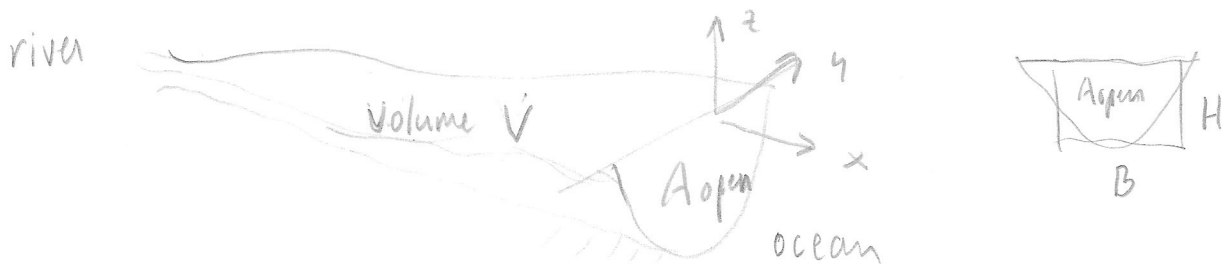
quintic



and recall $[u'] = U_E = \frac{g\beta\bar{s}_x H}{4g} \frac{H^2}{K} \sim \bar{s}_x$ (assume $A=K$) (2)

so $[s'] = \frac{H^2}{K} \bar{s}_x U_E \sim \bar{s}_x^2$ so stratification is very sensitive to \bar{s}_x

Next consider the volume integrated salt balance



$$\frac{d}{dt} \int_V s \, dV = - \int_{A_{open}} u s \, dA \approx - \int_{-B/2}^{B/2} \int_{-H}^0 (\bar{u} + u')(\bar{s} + s') \, dz \, dy$$

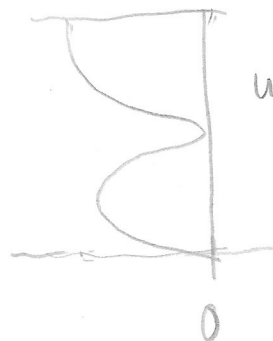
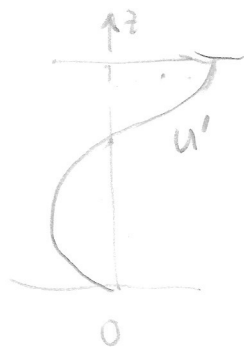
$$\approx HB \left(-\bar{u}\bar{s} - \overline{u's'} \right)$$

① ②

Most important terms are

① River flow $\times \bar{s}$ removes salt $(-\bar{u}\bar{s})$

② Exchange flow \times stratification brings salt in $(-\overline{u's'})$



$u's'$ always negative

$$\Rightarrow -\overline{u's'} > 0$$

and $-\overline{u's'} \propto \bar{s}_x^3 \star$

In steady state we can use the salt flux through the mouth

$$0 = -\underbrace{\bar{u}\bar{s}}_{(1)} - \underbrace{\overline{u's'}}_{(2)} \quad (*)$$

to come up with an equation for $\bar{s}(x)$.

... After some algebra we can write (*) as

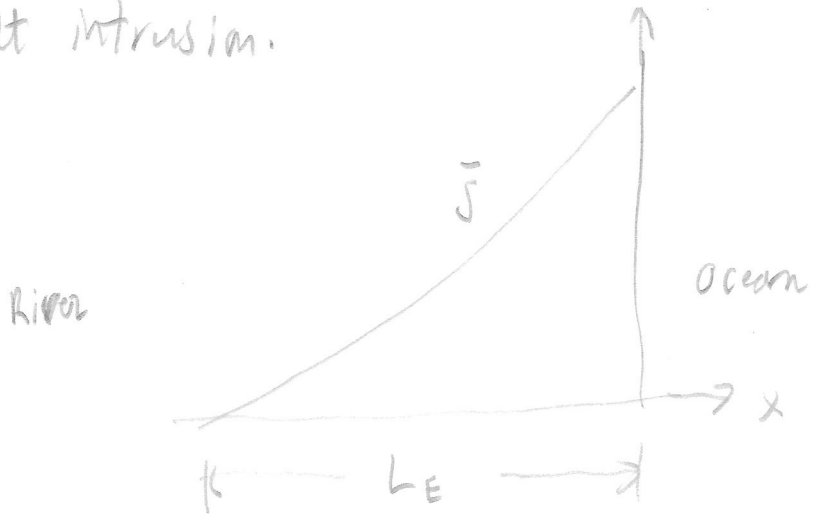
$$(1) \quad L_E^3 \sum_{(2)}^3 x - \sum_{(1)} = 0 \quad \text{where } \Sigma \equiv \bar{s}/s_{ocn}$$

$$\text{and } L_E = 0.024 c \left(\frac{\bar{u}}{c}\right)^{-1/3} \frac{H^2}{K}$$

and $c = \sqrt{g\beta s_{ocn} H} \sim 2 \times$ the fastest possible internal wave, with

$$\frac{\Delta\rho}{\rho_0} = \beta s_{ocn}$$

And L_E is the scale of the length of the salt intrusion.



Solving (i), guess $\Sigma = \alpha x^n$

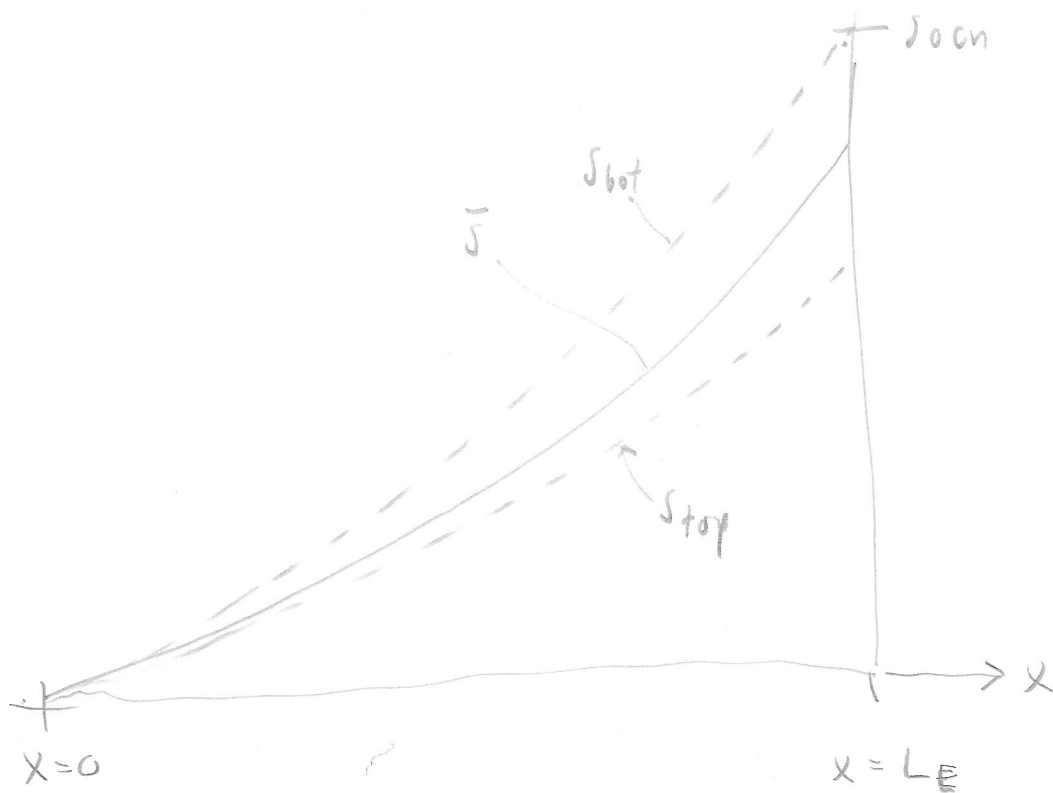
$$\Sigma_x = n\alpha x^{n-1}$$

$$(ii) \quad L_E^3 n^3 \alpha^3 x^{3n-3} - \alpha x^n = 0$$

In order for the solution to work for all x

$$\text{we require } 3n-3 = n \Rightarrow n = 3/2$$

$$\Rightarrow \Sigma = \alpha x^{3/2}$$

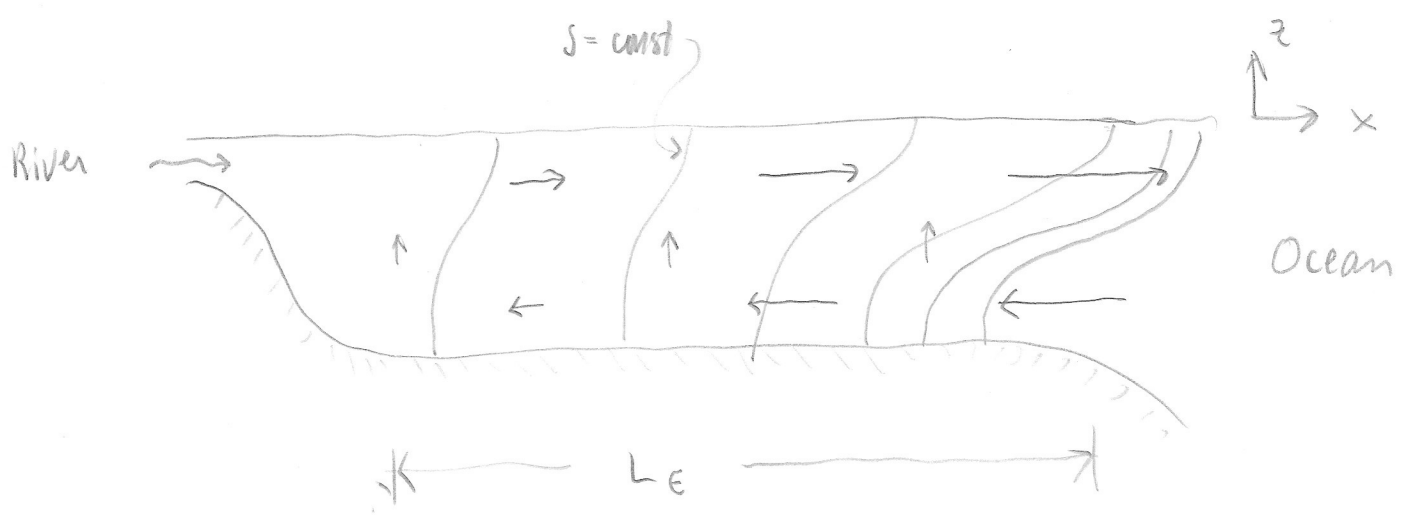


$$\text{So } \bar{J} \sim x^{3/2}$$

$$J' \sim x$$

$$u' \sim x^{1/2}$$

$$\bar{J}_x \sim \frac{J_0 c n}{L_E}$$



Scaling: Recall $L_E = 0.024 c \left(\frac{\bar{u}}{c}\right)^{-1/2} \frac{H^2}{K}$, $c = \sqrt{g\beta S_{00} H}$

and $\bar{J}_x \sim \frac{S_{00} c}{L_E}$

$$U_E = \frac{g\beta \bar{J}_x H}{48} \frac{H^2}{K} \sim \frac{c^2}{48} \frac{L_E}{L_E} \frac{H^2}{K}$$

$$\left[\frac{S'}{S_{00}}\right] \sim \frac{H^2}{K} \frac{\bar{J}_x}{S_{00}} U_E \sim \left(\frac{L_E}{L_E}\right)^2 \left(\frac{H^2}{K}\right)^2 c^2$$

6

Scaling

$$L_E \sim c \left(\frac{\bar{u}}{c} \right)^{-1/3} \frac{H^2}{K}$$

$$c = \sqrt{g\beta s_{ocn} t}$$

\Rightarrow L_E - shorter for more mixing ($\sim \frac{1}{K}$)

- shorter for more river flow

but it is stiff $L_E \sim Q_R^{-1/3}$

$$u_E \sim g\beta \bar{s}_x H \frac{H^2}{K} \sim \frac{g\beta s_{ocn} t}{L_E} \frac{H^2}{K} = \frac{c^2}{L_E} \frac{H^2}{K}$$

so $u_E \sim c \left(\frac{\bar{u}}{c} \right)^{1/3}$

$$\sim \frac{u_E}{c} \sim \left(\frac{\bar{u}}{c} \right)^{1/3}$$

- weak dependence on Q_R

- no dependence on K

(Why?)

(7)

Similarly we find

$$\frac{[S']}{\rho_0 c u} \sim \frac{U_E}{L_E} \frac{H^2}{K} \sim \frac{C^2}{L_E^2} \left(\frac{H^2}{K} \right)^2$$

so

$$\left[\frac{S'}{\rho_0 c u} \right] \sim \left(\frac{U}{C} \right)^{2/3}$$

again: independent of K